

Lost Creek Underground Storage Pilot – 2018 Groundwater Investigation Report

Prepared for:

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1 EXECUTIVE SUMMARY

Additional water storage is widely recognized as an important need for successful water resources management in Colorado's South Platte River Basin. There are many storage methods and locations to be considered. This report focuses on underground water storage (also called managed aquifer recharge [MAR], or aquifer storage and recovery [ASR]), in the unconsolidated alluvial aquifers of the South Platte River and its tributaries. More specifically, this study has evaluated field methods for assessing the feasibility of underground water storage in localized areas. Several regional underground water storage studies have been previously completed using desktop methods.

This study was a collaborative project between Aurora Water and Castle Rock Water. Additional THIS STUDY BUILDS ON PREVIOUS RECHARGE INVESTIGATIONS AND IS PRESENTED AS CURRENT BEST PRACTICES FOR FEASIBILITY EVALUATION OF UNDERGROUND WATER STORAGE IN THE ALLUVIAL AQUIFERS OF THE SOUTH PLATTE RIVER AND ITS TRIBUTARIES.

funding was provided by the Colorado Water Conservation Board. Leonard Rice Engineers, Inc. has worked with Aurora Water and Castle Rock Water to plan, implement, and document this project. The Lost Creek Ground Water Management District has participated as a stakeholder and has provided valuable input throughout the project.

Previous statewide and regional studies have ranked the Lost Creek Basin as a preferable location for underground storage due to deep bedrock, thick unsaturated alluvial deposits, and the unique water administration from being primarily within a Designated Basin¹. For these reasons, the Lost Creek Basin was chosen for this study. During the initial project planning, field investigation locations were identified where: 1) existing data indicated recharge was feasible; 2) unsaturated alluvial deposits are relatively thick; and, 3) property owners would allow field investigation on their property. Three locations were selected for field investigation but are not accurately located due to property owners' requests to remain anonymous.

We identified field investigation methods that were cost effective and have been used previously for alluvial groundwater investigations. Field investigation methods included geophysical surveys, surface infiltration tests, exploratory borehole drilling, and downhole infiltration testing. These methods were tested and compared to evaluate which are most appropriate for assessing local aquifer capacity for water infiltration. To some extent, the results from this investigation can be extrapolated for regional evaluation of water infiltration, storage, and recovery.

Table 1-1 below presents the various methods used and an assessment of their advantages and disadvantages.

¹ Designated Basins are areas in the Eastern Plains of Colorado where surface water is scarce and the primary supply is groundwater. Designated Basins are administered by the Colorado Groundwater Commission and local Groundwater Management Districts, outside of Colorado's prior appropriation water rights system.



| Field Method | Advantages | Disadvantages |
|---|--|--|
| Sonic Exploratory Borehole Drilling | Provides accurate and detailed continuous lithologic logs/samples Provides access for infiltration testing at depth Provides the highest quality lithology data that can be used to calibrate geophysical data | Relatively expensive Drilling disturbs the land surface and may interfere with other land uses Data limited to one location and requires multiple boreholes to investigate spatial heterogeneity |
| Time Domain Electromagnetics (TDEM) Geophysical Surveys | Interpreted results provide information about subsurface material grain size and layering Equipment can be configured for different resolution at different depths | Does not provide important cementation or consolidation information As configured for this project, TDEM did not provide adequate resolution to characterize the vertical heterogeneity observed in boring lithology 3-D interpretations were not confident as the results did not correlate with observed lithology |
| Frequency Domain Electromagnetic (FDEM) Geophysical Surveys | Results can be interpreted to represent bulk grain size data for upper 30 feet of the subsurface Interpretations were confident because they closely correlated to borehole lithology logs | Does not provide important cementation or consolidation information Provides a bulk measurement of the top part of the aquifer and therefore cannot provide important layering information |

Table 1-1 Field Investigation Methods Used During the Project



| Field Method | Advantages | Disadvantages |
|---|---|--|
| Electrical Resistivity Tomography (ERT) Geophysical Surveys | Interpreted results can relate electrical resistivity to subsurface material grain size | • Did not provide a fine enough resolution to characterize the vertical heterogeneity of the alluvial deposits |
| Multi-channel Analysis of Surface Waves (MASW) Geophysical Surveys | • Shear wave velocity can be interpreted to indicate density, and extrapolated to represent the degree of cementation and the depth to consolidated bedrock | Does not provide information about subsurface material grain size |
| Double-Ring Infiltrometer Surface Infiltration Testing | Provides a direct measurement of infiltration rates for surficial materials | Data limited to one location and requires multiple tests to investigate spatial heterogeneity Tests only the near-surface soils and the results are not influenced by deeper soil horizons |
| Constant-Head and Falling-Head Downhole Infiltration Testing | Provides a direct measurement of infiltration rates at depth Provides observations of relatively low and high permeability zones | Results are uncertain and interpreted saturated hydraulic conductivity values can vary significantly from true values Data limited to one location and requires multiple tests to investigate spatial heterogeneity |

1.1 FIELD METHOD CONCLUSIONS

We co-located field method testing sites so that results from different methods could be compared. The following is a summary of findings regarding the various field investigation methods tested:

• Surface geophysical surveys are a reliable means to characterize large areas of the subsurface, but should be locally calibrated and verified by exploratory boreholes. The geophysical results can provide valuable information for siting exploratory boreholes.



- The exploratory borehole lithology observations were the most detailed and accurate characterizations of subsurface materials, and were used as the baseline data to compare other results. The sonic drilling method provided high quality, continuous soil core samples necessary for characterizing the high degree of vertical heterogeneity of the lithology.
- FDEM interpreted results closely matched the exploratory borehole observed net sand thickness in the top 30 feet of the subsurface. FDEM results provide a reliable, relative characterization of the bulk grain size in the top 30 feet of unsaturated alluvial deposits.
- TDEM results did not have the resolution necessary to provide useful subsurface characterization for this project. The method could be adapted to have multiple passes with different equipment configurations (coil sizes). With proper configuration, TDEM may provide a reliable, relative characterization of the grain size and layering in unsaturated alluvial deposits.
- MASW interpreted results closely matched the borehole aquifer cementation and bedrock depth observations. MASW results provided a relative measure of cementation and bedrock depth in unsaturated alluvial deposits.
- ERT results were unreliable and are not recommended for future investigations.
- The interpretations of saturated hydraulic conductivity from the downhole infiltration tests are known to be uncertain. Therefore, the interpreted values are meant to give a relative, qualitative permeability estimate and not a precise value suitable for engineering design purposes.
- Surface double-ring infiltrometer testing provided a direct measurement of infiltration rates into surficial soils.
- The methods presented here are not exhaustive and additional field investigation technologies and methods should continue to be considered.

1.2 LOST CREEK BASIN HYDROGEOLOGIC CHARACTERIZATION CONCLUSIONS

We have made the following conclusions regarding the hydrogeologic characteristics of Lost Creek Basin as they relate to underground water storage planning:

- The alluvial deposits are vertically variable with frequent very thin to thick alternating beds of sand, clay, clayey sand, and sandy clay, with some gravelly areas. Intermediate clay beds were observed in all three test holes. These intermediate clay beds have the potential to present challenges for groundwater recharge. The lateral continuity of these beds and their potential effect on infiltration rates is important and should be further characterized.
- Near surface (shallower than 10 feet below ground surface) fine grained material may inhibit infiltration and should be specifically investigated, even in areas where the deeper deposits are permeable.



- Significant cementation of the alluvial aquifer vadose zone material was observed in two of three test holes.
- The degree of cementation influences permeability and porosity of the alluvial deposits and is a key planning consideration for underground storage at this location.
- Older alluvial deposits generally form topographic highs (terraces), are weakly to well cemented, and display a low to moderate permeability that is less conducive to infiltration. In these older alluvial deposits, the permeability of the subsurface material is not correlated to the gradation and is controlled by the cementation.
- Younger alluvium exists in topographically low spots, was generally loose, poorly cemented, and displayed moderate to high permeability; which is more conducive to infiltration.
- These conclusions are specific to the areas of the Lost Creek Basin that were investigated but are likely applicable to underground storage planning in other alluvial aquifers of the South Platte River and its tributaries.



1.3 RECOMMENDATIONS

Planning of underground storage will benefit from a phased feasibility investigation approach starting with regional or basin-scale investigations followed by localized field investigation. We recommend the following for future feasibility studies of underground storage in the alluvial aquifers of the South Platte River and its tributaries.

1.3.1 Basin Scale or Regional Investigations

- The cementation of the older alluvial deposits is a critical underground storage consideration and is not identified in regional studies, such as the 2011 Lost Creek Basin Aquifer Recharge and Storage Study (CGS, 2011). The reduction in pore space and permeability due to cementation may result in a lower amount of available storage than previously characterized. A desktop study should be conducted to re-evaluate available storage and permeability in the aquifer assuming that the older alluvial deposits are moderately to well cemented.
- Field mapping should be used to refine the available surficial geology and soil maps with a focus on mapping the younger and older alluvial deposits.
- Two-dimensional seismic profiles may help to characterize the distribution of cemented aquifer materials over larger areas.
- Basin/regional scale investigations and underground storage planning may also benefit from three-dimensional geologic modeling and groundwater flow modeling.
- If the conclusions of this study continue to hold true with additional investigation, local field investigations for recharge potential should be focused in the younger Piney Creek and Post Piney Creek alluvial deposits.

1.3.2 Local or Parcel Scale Investigations:

- The potential effect of shallow clay layers in the younger alluvium and the lateral continuity of these clay layers should be further investigated.
- Future investigations should include longer term infiltration testing using test pits or pilot infiltration basins to better characterize realistic infiltration rates. These tests should include observations of the wetting front and changes in groundwater levels throughout the test.
- Future parcel scale investigations should utilize FDEM geophysical surveys to optimize the location of infiltration basins and to target the relatively coarse materials in the top 30 feet of the alluvium.
- Future investigation utilizing the TDEM geophysical method should modify the equipment configuration to investigate the shallower materials in greater resolution.



- Future site investigations should include characterization of alluvium cementation. We recommend the use of seismic geophysical methods (such as MASW) in combination with test holes to characterize the degree of cementation in the alluvium.
- Future local-scale investigations should consider additional methods for investigating the existence and impact of vertical layering and lateral continuity of layers.
- Three dimensional geologic modeling and vadose zone/groundwater modeling should be conducted to evaluate the recharge and associated recovery of recharged water.

The results of this study continue to show that there is good potential for underground storage of water in the alluvial aquifers of the South Platte River and its tributaries. Using the recommended investigation methods presented herein will increase the success of underground storage projects.



2 INTRODUCTION

The City of Aurora (Aurora), in partnership with the Town of Castle Rock (Castle Rock), contracted Leonard Rice Engineers, Inc. (LRE) to evaluate field investigation methods related to assessing underground water storage potential. The objectives of the field investigations are to refine previous desktop hydrogeological characterization and to evaluate the effectiveness of subsurface investigation methods. The Lost Creek Ground Water Management District (LCGWMD) provided input on the project as an interested stakeholder.

Water storage has been identified as a critical need for the State of Colorado (State) to overcome the projected gap between future water supplies and demands. The 2016 State Water Plan identified 400,000 acre feet of water storage needs by 2050. The South Platte Basin has the majority of Colorado's water demand with most of the population and a significant portion of its agriculture. Identifying storage opportunities in the South Platte Basin is a crucial component of closing the State's future water supply gap. The South Platte Basin Implementation Plan (SPBIP, <u>http://southplattebasin.com</u>) calls for storage projects, including underground water storage or aquifer storage and recovery (ASR).

Following on the prior statewide and basin-scale planning efforts, the 2016 Colorado Legislature appropriated State funding for underground storage pilot projects in the Colorado Water Conservation Board (CWCB) Projects Bill (SB16-174). The legislative language specific to underground water storage is quoted below.

SECTION 8. Underground storage pilot project - appropriation. (1) For the 2016-17 state fiscal year, \$200,000 is appropriated to the department of natural resources for use by the Colorado Water Conservation Board. This appropriation is from the Colorado Water Conservation Board construction fund created in section 37-60-121, C.R.S. To implement this section, the Colorado Water Conservation Board may use this appropriation to conduct an underground storage pilot project to further evaluate the suitability of various aquifers to store water, availability of water to be stored, and a conceptual framework to initiate an underground storage project.

(2) The money appropriated in subsection (1) of this section remains available for the designated purposes until the project is completed.

The CWCB awarded SB16-174 grant funding to this project under Colorado Department of Natural Resources PO# POGG1 PDAA 201700001081.

LRE conducted site investigations at three locations in the Lost Creek Basin (Site A, Site B, and Site C). Time Domain Electromagnetics (TDEM) and Frequency Domain Electromagnetics (FDEM) investigations were conducted at all three locations. A surface electrical resistivity tomography (ERT) survey was conducted at two of the locations (Site A and Site C). One test hole was completed at Site A and two test holes were completed at Site B. Exploratory test holes were drilled with a sonic drill rig. Downhole permeability testing was conducted during test hole drilling and sampling. Multichannel Analysis of Surface Waves (MASW) seismic reflection surveys were conducted at each of the test hole locations. Finally, four surface infiltration tests were conducted at Site A and three surface infiltration tests were conducted at Site B (**Figure 1**).



3 LOST CREEK DESIGNATED BASIN GEOLOGY AND HYDROLOGY

A geologic map of the basin and summary of the geologic units of the Lost Creek Basin are presented in **Appendix A**. This is the best available geologic map of the area, and to our knowledge, the Lost Creek Basin has not been mapped in greater detail.

The basin is composed of the following geologic deposits, listed from youngest to oldest (Colorado Geologic Survey [CGS], 2011):

- Holocene and Pleistocene age eolian sands and loess up to 100 feet thick. The deposits form long, narrow, dunes trending northeast southwest.
- Holocene age Piney Creek and Post-Piney Creek Alluvium is located along the current low lying drainages and dry river channels. Throughout this report the Piney Creek and Post-Piney Creek Alluvium is often referred to as "younger alluvium". The younger alluvium is composed of clay, silt, sand, gravel, and cobbles, and has less cementation.
- Pleistocene age alluvial deposits are typically located above the existing drainage and active river channels in the basin. Throughout this report these Pleistocene alluvial deposits are referred to as "older alluvium". The older alluvial deposits are composed of clay, silt, sand, and gravel, are more cemented, and are often roughly to well stratified with distinct layering.
- The Quaternary units are underlain by Cretaceous and Tertiary sandstone, shale, and claystone of the Denver Basin (Denver, Arapahoe, Laramie, and Fox Hills units).

The Lost Creek Basin is an ephemeral tributary of the South Platte River. The main Lost Creek channel is fed from the south by its ephemeral tributaries Long Draw and Sand Creek. The streams do not maintain flow, are not fed directly by groundwater and only flow in response to precipitation events. The total drainage area is approximately 433 mi², approximately 43 miles long and 14 miles wide (CGS, 2011).

A basin-wide potentiometric surface was interpreted in CGS, 2011, based on recorded well water levels. In general, water levels are deeper in the southern and central parts of the basin (50 feet to greater than 100 feet below ground surface) and are shallower to the north, closer to the South Platte River (as high as 10 feet below ground surface). The slope of the water table results in the maximum unsaturated alluvium thickness in the southern and central portion of the aquifer, with less unsaturated alluvium in the northern part of the basin (CGS, 2011).



4 PREVIOUS STUDIES

The State has funded several previous studies that evaluated the feasibility of underground water storage at the State and basin scale (**Table 4-1**).

| Title | Publisher | Author | Year |
|--------------------------|--------------------------------|--------------------------|------|
| Artificial Recharge of | Colorado Geological Survey | Ralf Topper, Peter E. | 2004 |
| Ground Water in Colorado | | Barkmann, David A. Bird, | |
| - A Statewide Assessment | | and Matthew A. Sares | |
| SB06-193 Underground | Colorado Water Conservation | CDM | 2007 |
| Water Storage Study | Board | | |
| Lost Creek Basin Aquifer | Lost Creek Ground Water | Nicholas Watterson and | 2011 |
| Recharge and Storage | Management District | Ralf Topper | |
| Study | and Colorado Geological Survey | | |
| | | | |
| South Platte Storage | South Platte Basin Roundtable | Stantec and Leonard Rice | 2017 |
| Study | | Engineers | |

5 SITE INVESTIGATIONS

LRE conducted site investigations at three locations within the southern portion of the Lost Creek Basin. During the initial project planning, field investigation locations were identified where: 1) existing data indicated recharge was feasible; 2) unsaturated alluvial deposits are relatively thick; and, 3) property owners would allow field investigation on their property. Three locations were selected for field investigations. Sites A and B were on private property and the locations of which are not presented in this report at the property owner's requests. Site C is located along a public right of way and the location is included in this report.

The field investigation was conducted in a phased manner to utilize the data collected in each phase to inform the investigations of the following phases. Geophysical investigations were initially conducted at all three sites. Based on the preliminary results of the geophysical investigations, one test hole location was selected at Site A and two test holes locations were selected at Site B. Additionally, four surface infiltration test locations were selected at Site A and three surface infiltration test locations were selected at Site B. **Figure 1** and **Figure 2** shows the field investigation locations for each site.

5.1 GEOPHYSICAL INVESTIGATIONS

Olson Engineering (Olson) was subcontracted by LRE to perform all geophysical surveys associated with the project. The types of geophysical surveys conducted measure variations in electrical



resistivity (or its reciprocal, conductivity) and seismic wave travel times. The collected data are then interpreted, and correlated with other lines of evidence, to assess subsurface properties such as changes in lithology, cementation, moisture content, and the relative amount of coarse or fine grained material. The objectives of the geophysical investigation were to:

- Evaluate the vertical layering in the Lost Creek Basin,
- Investigate potential for laterally continuous clay beds,
- Characterize the bedrock contact, and
- Aid in siting test hole and surface infiltration locations.

The following descriptions of the geophysical techniques used were provided to LRE by Olson.

5.1.1 Time Domain Electromagnetics (TDEM) [Provided by Olson]

Olson used the Dynamic NanoTEM (DNT) system by Zonge International, Inc. to acquire the time domain electromagnetic data. The DNT allows for continuous acquisition of TDEM data and incorporates positional data via an integrated GPS. The TDEM survey technique involves transmitting a 50% duty cycle, time domain, square-wave signal into an ungrounded loop of wire. This square wave signal alternates between positive, zero, negative, and zero voltages. The measurements are made when the transmitter is off, i.e., when the transmitter voltage is zero. During these times, decaying magnetic fields from subsurface conductors can be measured. The decaying voltages can be mathematically modeled, providing a vertical sounding of resistivity beneath the receiver loop.

The DNT transmitter is capable of turning off in less than 2 microseconds, and the receiver's first measurement point (corresponding to the shallowest data) is at 4.5 microseconds after turn-off. This fast turn-off and high sample rate allow for good vertical resolution in the near-surface. The turn-off time and sample rate can be adjusted according to the size of the transmitter loop, where the size of the loop is designed specifically for the depth of investigation and the resolution required.

Depth of investigation is dependent on the size of the wire loops while lateral resolution is controlled by the down line station separation. In the field, two square loops of wire were used for the transmitter and receiver loops. This is called a coincident loop as the transmitter and receiver wire are coincident with each other. The signal is transmitted into the transmitter loop and received by the receiver loop. The transmitter loop was 40 x 40 feet in Site A and Site B, and 15 x 65 feet in Site C. Data were collected over Site A and Site B at line spacings of 196 to 295 feet and 330 to 560 feet, respectively, with both north-south and east-west traverses, and over Site C a single 2,590 feet line was collected in the east-west direction.

The advantage of the DNT is the ability to continuously acquire TDEM data while moving the transmitter-receiver loop along the survey profile. Data acquisition with the DNT allows for dense down-line station spacing. Data were continuously recorded at 32 Hz and then stacked during data processing, resulting in average down-line station spacing of less than 3 feet. The receiver operator monitored data quality and positional data in real time to verify system performance. The DNT system has an integrated Global Positioning System (GPS) capable of decimeter level accuracy for



sensor positioning. GPS data were recorded at a rate of one reading every second and were written directly into the raw data file.

The results of the TDEM survey at the three survey locations are presented on **Figure 3** through **Figure 8**.

5.1.2 Frequency Domain Electromagnetics (FDEM) [Provided by Olson]

An FDEM instrument consists of a transmitter coil which generates a primary electromagnetic field at a specified frequency separated by some distance from a receiver coil which measures the response (secondary) electromagnetic field generated due to the interaction of the soils to the primary field. This allows for simultaneous measurements of both the in-phase and the quadrature components of the secondary magnetic field generated in the receiver coil. The quadrature component is most sensitive to electrical conductivity, generally due to changes in lithology, moisture, and/or fines (clay) content, whereas the in-phase component is most sensitive to magnetic susceptibility, generally due to the presence of metallic features in the subsurface, either ferrous or non-ferrous.

The FDEM data were acquired using a DUALEM-4 terrain conductivity meter mounted to a cart and were recorded at a rate of 1 Hz, using a primary field of 9.8 kHz with an antenna spacing of 12 feet. GPS measurements were embedded into the data stream at a rate of 1 Hz to allow for precise data positioning. Data were collected over Site A and B at line spacing of 260 to 390 feet, with both north-south and east-west traverses. Data were collected over Site C in a single 2,790 feet line spacing in the east-west direction.

The DUALEM is generally sensitive to a depth range of approximately 10 to 20 feet. This sensitivity range, or the "effective depth", is a response to the orientation and spacing of the transmitter and receiver coils, the frequency of the primary field, and the bulk electromagnetic properties of the subsurface. It is important to understand that each FDEM measurement is effectively a 'single-depth' response; therefore, FDEM results are indicative only of lateral subsurface variations.

The results of the FDEM survey at the three survey locations are presented on **Figure 9** through **Figure 11**.

5.1.3 Electrical Resistivity Tomography (ERT) [Provided by Olson]

In an ERT survey, an electrical current is injected into the ground through two electrodes. Voltages on the surface are measured using another pair of electrodes revealing the direction and amount of current flow in the subsurface. These data are interpreted in terms of the resistivity (bulk resistance to current flow) of the earth materials.

ERT data at this site were acquired using a dipole-dipole electrode array configuration for mapping relatively shallow lateral variability with high resolution. The data were collected using an IRIS SYSCAL Pro system. For each line section, 72 to 96 electrodes were deployed at 16 foot intervals for a total of 1200 feet (Site A) and 1500 feet (Site C), respectively. A 12-volt external battery provided the current transmitted into the electrode arrays.

The results of the ERT survey at the two survey locations are presented on **Figure 12**, and **Figure 13**.



5.1.4 Multi-Channel Analysis of Surface Waves (MASW) [Provided by Olson]

Active and passive MASW surveys were conducted at each study area to analyze the velocities of surface waves and shear waves. The observed velocities are indicators of the degree of hardness of the subsurface material (i.e. bedrock or cemented alluvium).

Using a Geometrics Geode 24-channel seismograph with 24 4.5 Hz vertical component geophones placed at 10 foot intervals, MASW data were acquired using an active seismic source consisting of a sledgehammer impacting a plastic strike plate. Shot points were located every 30 feet, beginning with a 30 foot off-end shot at the beginning of each line. Acquisition parameters of the seismic system for the active MASW methods comprised of stacked 2 second records at a 0.125 millisecond (ms) sample rate. Passive MASW data were acquired as well, using acquisition parameters comprised of fifteen 30-second records with a 2 ms sample rate. There are no predefined source points for passive-source surface seismic surveys. Instead, the method uses ambient noise, or vibrational energy, that exists at a site. A sledgehammer impacting a plastic plate off the end of the line was used to add extra signal to the passive measurement to ensure maximum depth of penetration.

The results of the MASW survey are plotted on the test hole logs included in **Appendix B** and are summarized on **Figure 14**.

5.2 SURFACE INFILTRATION TESTING

LRE selected 7 locations to conduct surface infiltration testing based on the results of the TDEM and FDEM survey. The purpose of the infiltration testing was to evaluate the permeability of the near surface soils. The infiltration testing targeted both electrically resistive (coarse grained) and conductive (fine grained) areas to attempt to correlate infiltration rates to electrical resistivity. Infiltration tests were conducted using double ring infiltrometers and following ASTM D3385-18. The tests were run for a duration of 4-hours or until a steady state infiltration rate was achieved. A summary of the infiltration results is presented in **Table 5-1**.

The steady state infiltration rates from double ring infiltrometers are not equivalent to the saturated hydraulic conductivity (k_s) of the surficial soils unless steady state infiltration with the water table is achieved. Because the water table at the sites is approximately 80 feet below ground surface (or deeper), steady state infiltration with the water table was not achieved. Therefore, the infiltration rates provided in table 5-1 are greater than that of the material's k_s .



| Material Tested | Infiltration Test Number | Steady State Infiltration Rate (ft/day) | Geometric Mean - Steady State Infiltration Rate (ft/day) | Arithmetic Mean - Steady State Infiltration Rate (ft/day) | |
|---|-----------------------------|---|---|---|--|
| Older | I-1 | 0.6 | | | |
| Alluvium Terraces | I-2 | 2.1 | 2.0 | 3 | |
| | I-3 | 6.3 | - | | |
| | I-4* | 0.3 | | 5.3 | |
| Younger | I-5 | 13.5 | 2.7 | | |
| Alluvium | I-8 | 3.6 | | | |
| | I-9 | 3.8 | | | |
| *Near surface deposits at I-4 are clays and the soil gradation likely dominated the low permeability test result. | | | | | |

Table 5-1 - Surface Infiltration Test Results

5.3 TEST HOLE DRILLING AND DOWNHOLE PERMEABILITY TESTING

LRE selected 3 locations to advance test holes and conduct downhole permeability testing. The purpose of the test holes were to: (1) characterize the lithology, (2) conduct downhole permeability testing to characterize the unsaturated aquifer material, and (3) provide calibration data for the geophysical surveys. Lithology logs and permeability testing summaries of the test holes are included in **Appendix B**. The test hole locations are shown on **Figure 1**.

5.3.1 Drilling and Lithological Logging

Test holes were drilled via the sonic method. The sonic drilling method consists of advancing an inner core barrel and an outer drive casing to obtain a continuous core sample of the formation (inner core barrel) while maintaining the open borehole (outer drive casing). The drill rods are advanced using a high-frequency resonant energy that vibrates the casing and causes liquefaction of the unconsolidated material encountered. This method was chosen because it provides for continuous and relatively undisturbed sampling of the test hole typically with 100% sample recovery. This allows for detailed lithological logging of the entire borehole. A summary of the lithology observed during drilling is presented in **Table 5-2**.



| Test hole | Total Depth | Static Water Table | Depth to Bedrock | Lithology |
|--------------|----------------|-----------------------|---------------------|---|
| 0-1 | 80 feet | Not encountered | Not encountered | Loose to well cemented sand and clay Dry to depth Frequent calcium carbonate precipitate Varying degrees of cementation with some samples having the strength and appearance of weak sandstone Massive bedding to very thin bedding (distinct horizontal structure) at different depths Drilling was more difficult than typical for unconsolidated sands and clays |
| 0-4 | 60 feet | Not encountered | Not encountered | Same as 0-1 except with generally higher cementation Driller noted that the rig response was typical to sandstone and other consolidated deposits |
| 0-5 | 85 feet | 69 feet | 81 feet | Loose to weakly cemented gravel, sand and clay Shallow samples were moist, moisture decreased with depth until static water level was reached Typically massively bedded with infrequent distinct horizontal structure (very thinly to thinly bedded) Zones with a distinct horizontal structure generally correlate with zones of weak cementation Blue shale and sandstone (Denver Formation) encountered at 81 feet. |



5.3.2 Downhole Permeability Testing

Downhole permeability testing was conducted using constant-head and falling-head test methods. The following steps were followed during permeability testing:

- 1. The outer drive casing (6-inch diameter) was advanced to the top of the desired test interval.
- 2. The inner core barrel casing (4-inch diameter) was advanced to the bottom of the test interval approximately 5 to 10 feet below the outer casing. The soil core and inner casing were then removed.
- 3. A perforated PVC access tube was advanced to near the bottom of the test hole.
- 4. A data logging pressure transducer was then placed inside the PVC access tube.
- 5. Water was pumped into the test hole.
- 6. The inflow rate was modified to maintain a constant head at a level as close as possible to the top of the test interval.
- 7. Metered water supply was poured into the feed pipe until the water level and flow rate stabilized. The flow rate required to maintain a constant-head within the test interval was recorded and used for interval permeability calculations.
- 8. Once the constant-head test was completed, the water level decline rate (falling head) was monitored for use in interval permeability calculations.
- 9. If a constant head could not be maintained at the minimum pump/meter flow rate (less than approximately one to two gallon per minute), the test was switched to the falling head method.
- 10. During a falling-head test, the casing was filled with water to a depth of approximately 10 feet above the test interval and the decline in head was monitored for at least 10 minutes.

The hydraulic conductivity of the subsurface material was then calculated from the constant- head and falling-head test data using methodology presented in the Bureau of Reclamation 2001 Engineering Geology Field Manual, Chapter 17 (USBR, 2001).

For the constant-head test, flow rate and water level were continuously monitored until both stabilized. The "Gravity Permeability Test Method 1" (USBR, 2001) was used to determine the saturated hydraulic conductivity within intervals where a stable flow rate and water level were observed. The equation is as follows:

$$K = \frac{2Q}{(C_s + 4)r T_u}$$

```
Where:

K = Hydraulic \ conductivity \ (ft/s)

Q = Flow \ rate \ into \ the \ test \ hole \ (ft^2/s)

C_s = Conductivity \ coefficient

r = Radius \ of \ the \ test \ section \ (ft)

T_u = Distance \ from \ the \ water \ surface \ in \ the \ test \ interval \ to \ the \ static \ water \ level \ (ft)
```



For the falling-head test, the water level decline rate was continuously monitored and the saturated hydraulic conductivity was calculated using the "Falling-Head Test Method for tests above the static water table" (USBR, 2001). The equation is as follows:

$$K = \frac{r_1^2}{2l\Delta t} \left[\frac{\sinh^{-1}\left(\frac{l}{r_e}\right)}{2} x \ln\left(\frac{2H_1 - l}{2H_2 - l}\right) - \ln\left(\frac{2H_1H_2 - lH_2}{2H_1H_2 - lH_1}\right) \right]$$

Where:

$$\begin{split} &K = Hydraulic \ conductivity \ (ft/s) \\ &r_1 = Inside \ radius \ of \ the \ drop \ pipe \ (ft) \\ &r_e = Effective \ radius \ of \ the \ test \ section \ (ft) \\ &\Delta t = Time \ intervals \ (seconds) \\ &l = Length \ of \ the \ test \ section \ (ft) \\ &sinh^{-1} = Inverse \ hyperbolic \ sine \\ &ln = Natural \ Logarithm \\ &H = Length \ of \ water \ column \ from \ bottom \ of \ test \ interval \ to \ water \ surface \ in \ standpipe \ (ft) \ at \ time \\ &of \ measurements \end{split}$$

The average calculated saturated hydraulic conductivity from each interval of the three test holes is presented in **Table 5-3**.

6 DATA INTERPRETATION AND EVALUATION OF INVESTIGATION METHODS

The results of the field investigation generally correlate with the delineation between the younger alluvium (Holocene age Piney Creek and Post-Piney Creek Alluvium) and older alluvium (Pleistocene age). See page 13 and **Appendix A** for details on the general geologic background of these units. Test holes 0-1 and 0-4 were drilled on topographic highs and in the older alluvium. In general, the subsurface materials observed at these locations were weakly to well cemented, dry, and were interpreted to have a moderate to low permeability. Test hole 0-5 was drilled in a topographic low and through the younger alluvium. The 0-5 subsurface materials observed were loose to very weakly cemented, moist, and interpreted to have a moderate to high permeability.

6.1 SONIC TEST HOLE DRILLING RESULTS

The layering of all the alluvial deposits encountered is complex with frequently alternating beds of sand, clay, sandy clay, and clayey sand. The complex layering was characterized in continuous core samples from the sonic drilling method. The continuous sampling allowed for detailed logging of the lithology and the identification of thin bedding. However, the thin bedding and frequently alternating beds was not captured in the geophysical surveys.



| | | | Saturated Hydraulic | | |
|-----------|-----------------|--------------------|--------------------------|------------------|---|
| Test hole | Top (ft bgs) | Bottom (ft bgs) | Conductivity (ft/dav) | Method | Notos |
| 0-1 | 10 | 20.5 | 0.6 | Falling | Could not maintain a constant head within the |
| | | | | Head | flow rate limitations. |
| | 20 | 30.2 | 11 | Constant Head | |
| | 30 | 40 | 11 | Constant Head | |
| | 40 | 49.4 | 0.1 | Falling Head | Could not maintain a constant head within the flow rate limitations. |
| | 50 | 60.5 | 7 | Constant Head | |
| | 59.7 | 70.5 | 0.2 | Falling Head | Could not maintain a constant head within the flow rate limitations. |
| | 69.7 | 80.5 | 50 | Constant Head | Potentially hit water near the base of the borehole, at approximately 79 - 80 feet. |
| | Geome | tric Mean | 2.4 | ft/day | |
| 0-4 | 9.9 | 20.5 | 0.045 | Falling Head | Could not maintain a constant head within the flow rate limitations. |
| | 19.9 | 30.2 | 10 | Constant Head | |
| | 29.8 | 40 | 0.02 | Falling Head | Could not maintain a constant head within the flow rate limitations. |
| | 39.8 | 50.5 | 0.009 | Falling Head | Could not maintain a constant head within the flow rate limitations. |
| | 49.8 | 60.5 | 0.081 | Falling Head | Could not maintain a constant head within the flow rate limitations. |
| | Geome | tric Mean | 0.09 | ft/day | |
| O-5 | 9.8 | 20.5 | 0.3 | Falling Head | Could not maintain a constant head within the flow rate limitations. |
| | 19.8 | 30 | 17 | Constant Head | |
| | 29.8 | 40 | 23 | Constant Head | |
| | 39.8 | 50 | 8 | Constant Head | |
| | 50 | 60 | 14 | Constant Head | |
| | 59.8 | 70 | 235 | Constant Head | Water table at 66-69 feet below ground surface. Only test in saturated zone |
| | Geome | tric Mean | 12.1 | ft/day | |

Table 5-1 - Downhole Permeability Testing Results and Interpretation



6.2 ELECTROMAGNETIC SURVEY RESULTS

6.2.1 Time Domain Electromagnetic Survey Results

The TDEM survey was used to record variations in electrical resistivity/conductivity data (i.e. due to changes in lithology and/or moisture) with depth to assess vertical layering. As shown in the composite lithology logs (**Appendix B**), there is no apparent correlation between the fine-grained beds and electrically conductive zones, or the coarse grained beds and resistive zones. The likely cause of this mismatch is the large 40-foot loops used during the TDEM survey. The larger loops allow for deeper depths of investigation but reduce the vertical resolution. Future TDEM surveyors should consider using a smaller loop-size to increase data resolution to characterize the vertical heterogeneity of the shallow alluvial deposits.

KEY TAKEAWAY

FUTURE TDEM SURVEYORS SHOULD CONSIDER USING A SMALLER LOOP-SIZE TO INCREASE DATA RESOLUTION TO CHARACTERIZE THE VERTICAL HETEROGENEITY OF THE SHALLOW ALLUVIAL DEPOSITS.

6.2.2 Frequency Domain Electromagnetic Survey Results

FDEM surveys were also conducted to record changes in electrical resistivity at relatively shallow depths (effective depths of up to 30 feet). The results of the FDEM survey correlate well with the thickness of sand observed in the top 10 and top 30 feet at the test holes. **Table 6-1** shows a comparison of the logged feet of sand and the FDEM Resistivity. Test hole 0-1 had the greatest thickness of logged sand and the highest electrical resistivity. 0-4 had less sand and a lower resistivity and test hole 0-5 had the least amount of sand and correspondingly the lowest measured electrical resistivity.

KEY TAKEAWAY

CEMENTATION OF THE ALLUVIUM APPEARS TO HAVE NO INFLUENCE ON THE FDEM RESULTS. FDEM appears to be an effective method for locating zones of relatively coarse and fine material in the top 30 feet of the subsurface. This could be an effective tool for relative comparison when siting the best location for an infiltration basin in a target parcel. It is important to note that cementation of the alluvium appears to have no influence on the FDEM results and FDEM should only be used to target beneficial infiltration locations if the degree of cementation of the particles has been confirmed with another method.



| Test hole | Feet of Sand ¹ in Top 30 Feet | Feet of Sand ¹ in Top 10 Feet | FDEM Resistivity (Ohm-m) |
|------------------|---|---|-----------------------------|
| 0-1 | 17.5 | 9.5 | 24.9 |
| O-4 | 15 | 5 | 15.1 |
| O-5 ² | 13 | 0 | 11.6 |

Table 6-2 - Comparison of FDEM Results and Observed Lithology

1. Defined as material logged as sand with less than 12% fines (SW, SP, SW-SC, SP-SC). Clayey sands (SC) were not counted.

2. Test hole O-5 was off the path of the FDEM and the results are from a nearby location.

6.2.3 Effects of Moisture Content/ Saturation on TDEM and FDEM Results

The degree of soil saturation does not appear to have had a significant impact on the TDEM and FDEM resistivity values. In general, both the TDEM and FDEM indicated more electrically-resistive materials (i.e. coarse materials) in the topographic lows and more conductive materials (i.e. fine material) near the topographic highs. The trend is not perfect with some variations and areas of more conductive material in the topographic highs. As we observed in test hole 0-5, the materials in the topographic lows appear to have a higher degree of saturation while the materials on the topographic highs were dry with no moisture. If the saturation was greatly influencing the conductivity of

KEY TAKEAWAY

THE DEGREE OF SOIL SATURATION DOES NOT APPEAR TO HAVE HAD A SIGNIFICANT IMPACT ON THE TDEM AND FDEM RESISTIVITY VALUES.

the materials, then it is expected that the materials in the topographic lows would be more conductive when they were actually observed to be more resistive. This indicates that the material gradation (i.e. coarse versus fine) is driving the observed electrical resistivity/conductivity.

6.3 ELECTRICAL RESISTIVITY TOMOGRAPHY RESULTS

The ERT survey was conducted to measure the electrical resistivity/conductivity of a relatively shallow vertical section, but with higher resolution than the TDEM survey. However, the ERT survey did not appear to produce significantly more detailed results than the TDEM survey. The fine layering of the alluvial material was not captured in the ERT results, which indicated relatively massive deposits that were not observed in the exploratory boreholes.

KEY TAKEAWAY

THE FINE LAYERING OF THE ALLUVIAL MATERIAL WAS NOT CAPTURED IN THE ERT RESULTS.



6.4 MULTI-CHANNEL ANALYSIS OF SURFACE WAVES (MASW) RESULTS

The effect of cementation was most clearly picked up with the MASW survey. As seen in the composite lithology logs, the areas of relatively high shear wave velocity tend to correlate with zones of observed cementation and high dry strength. **Figure 14** shows the shear wave velocities in profile at the three test holes. The shear wave velocities appear to be higher at 0-1 and 0-4 relative to 0-5, particularly between the depths of 20 to 60 feet below ground surface. The average shear wave velocities at 0-1 and 0-4 whre 905 and 847 ft/s, respectively. These values are significantly greater than the average shear wave velocity of 808 ft/s at 0-5. The relatively higher shear wave velocities of the cemented materials make MASW (or other seismic methods) an appropriate tool for delineating cemented versus loose materials. The cemented, older alluvial deposits had, on average, a shear wave velocity

KEY TAKEAWAY

THE HIGHER SHEAR WAVE VELOCITIES OF THE CEMENTED MATERIALS MAKE MASW (OR OTHER SEISMIC METHODS) AN APPROPRIATE TOOL FOR DELINEATING CEMENTED VERSUS LOOSE MATERIALS.

MASW COULD ALSO BE AN APPROPRIATE METHOD FOR ESTIMATING BEDROCK DEPTH BECAUSE OF ITS DENSITY AND RESULTING HIGH SHEAR WAVE VELOCITIES. 40 to 100 ft/s greater than the relatively loose deposits.

MASW could also be an appropriate method for estimating bedrock depth because of its relatively high density (consolidation and cementation). O-5 is the only test hole that encountered bedrock, where it was observed at 80 ft bgs. This correlates well with the MASW results that increase to the test hole maximum value of greater than 1,500 ft/s at a depth of approximately 78 feet.

The one-dimensional MASW results indicate that two-dimensional seismic profiles may provide appropriate characterization of the distribution of cemented aquifer material and bedrock depth across larger areas.

The MASW data did not show significant correlation to the grain size of materials observed in test hole samples.

6.5 DOWNHOLE PERMEABILITY TESTING

The interpretation of saturated hydraulic conductivity from the infiltration tests is known to be uncertain. The interpreted values are meant to give a relatively high or relatively low permeability result and the exact value is not suitable for engineering design purposes. Stephens, 1979 indicates that the USBR method may result in k_s values that vary up to 160% from the true value.

Additionally, the interpreted k_s from the falling head test is correlated to the driving head, which can be controlled by the amount of water above the testing area. The interpreted k_s values from the-falling head tests were often up to an order of magnitude lower than the interpreted k_s values from a constant-head test in the same interval. Despite this uncertainty, the USBR method seems to be the best available method for interpreting borehole permeability results in the unsaturated alluvium.



The main uncertainty in permeability or infiltration testing is the characterization of the wetting front. Characterization of a wetting front would allow for the application of more complex vadose zone numerical models and a more accurate interpretation of k_s . However, in downhole permeability testing, characterization of the wetting front is difficult and therefore, the USBR method

KEY TAKEAWAY

DOWNHOLE PERMEABILITY TESTING IS AN EFFECTIVE TOOL FOR QUALITATIVE CHARACTERIZATION OF SUBSURFACE MATERIAL PERMEABILITY.

is still the best tool available. Despite the method uncertainties, the downhole permeability testing is an effective tool for qualitative characterization of subsurface material permeability and is a good tool for comparing the relative permeability of alluvial layers. In the areas investigated for this project, the distinct differences in permeability due to cementation were evident in the permeability testing results; and the borehole permeability testing was an effective tool for broadly characterizing the effect of cementation on permeability.

6.6 SURFACE INFILTRATION TESTING

The surface infiltration tests provide an estimate of potential infiltration rates in the shallow subsurface. The observed infiltration rates are likely to vary from actual infiltration rates in a constructed basin due to the effects of deeper soil horizons, long-term operations achieving steady conditions with the groundwater table, lateral heterogeneity and clogging over time, none of which are considered in the double-ring infiltrometer tests. A long-term infiltration test in a test basin is a more suitable method for estimating long-term infiltration rates and basin design. The double-ring infiltrometer tests are an effective test for targeting soils with a relatively high infiltration capacity.

Saturated hydraulic conductivity may be estimated from the infiltration rates, but there is uncertainty in these estimates due to the effects of unsaturated soil physics on the infiltration rates. These effects are difficult to accurately account for without a wellcharacterized wetting front and understanding of the subsurface anisotropy. Therefore, the tests are best used for approximating relative infiltration rates instead of estimating absolute k_s values.

KEY TAKEAWAY

A LONG-TERM INFILTRATION TEST IN A TEST BASIN IS A MORE SUITABLE METHOD FOR ESTIMATING LONG TERM INFILTRATION RATES AND BASIN DESIGN.

6.7 ALLUVIAL AQUIFER CEMENTATION

The young and older alluviums were not significantly different with regard to material gradation with both being composed of interlayered sand and clay with some gravel. The distinct difference between the two ages of alluvium is the degree of cementation and the horizontal thinly bedded structure of the cemented units. Calcium carbonate cement has the potential to fill the void space of the alluvium resulting in a lower effective porosity and lower water storage capacity. The cementation also results in lower permeability, which reduces the rate at which the aquifer could be recharged through



surface infiltration basins. A reduction in permeability of the cemented units is evident in the downhole permeability testing data. The geometric mean of the interpreted permeability at 0-1 and 0-4 is 20% and 1% of that at 0-5, respectively.

For future siting of infiltration basins, it is important to know whether the cementation or the alluvium composition (i.e. sand versus clay content) has a greater impact on the material permeability. In general, coarser deposits have a high permeability and fine deposits have a low permeability. Therefore, it would be anticipated that the interpreted hydraulic conductivity from the permeability tests would correlate to the thickness of sand deposits in the 10-foot test interval.

A correlation between the interpreted hydraulic conductivity from each permeability test and the thickness of logged sand in the 10 foot test interval is presented on Figure 15. Each data point representative of one of the is permeability tests conducted (see Table 5-3). The lines are the logarithmic best fit for each set of permeability tests grouped by test hole. The correlation coefficient r-squared (R², unitless) for the permeability tests at test hole 0-4 is 0.0023. This is considered a poor fit, therefore the permeability of the material does not appear to be correlated to the gradation or relative coarseness of the alluvium. The R² of the permeability tests at test hole 0-1 is 0.1008 which

KEY TAKEAWAY

THE WEAK CORRELATION BETWEEN GRADATION AND PERMEABILITY IN THE OLDER ALLUVIUM INDICATES THAT THE PERMEABILITY OF THESE UNITS IS PRIMARILY CONTROLLED BY THE DEGREE OF CEMENTATION OF THE DEPOSITS. IN THE YOUNGER ALLUVIUM, WHERE THE ALLUVIUM IS NOT CEMENTED, THE PERMEABILITY IS CONTROLLED BY THE GRADATION OF THE DEPOSITS.

indicates a weak correlation of the permeability to the gradation or relative coarseness of the alluvium. Finally, the R² of the permeability tests at 0-5 is 0.5971 which indicates that there is a statistical correlation between the coarseness of the deposit and the permeability. The weak correlation between gradation and permeability in the older alluvium indicates that the permeability of these units is primarily controlled by the degree of cementation of the deposits. In the younger alluvium, where the alluvium is not cemented, the permeability is controlled by the gradation of the deposits.

6.8 AQUIFER LAYERING

As observed in the lithologic logs, complex layering and frequently alternating beds of sand, clay, sandy clay, and clayey sand are present in each of the test holes. The effect of the fine structure on aquifer recharge is uncertain at this time. The lateral continuity of thin clay beds could not be determined with the data collected. A continuous thin clay bed could reduce the rate of groundwater recharge. If the clay beds are inter-fingered or spatially limited, sufficient pathways for recharge water would be present and the observed thin beds would have limited effect on the ability to recharge the aquifer with infiltration basins.



7 CONCLUSIONS AND RECOMMENDATIONS

7.1 FIELD METHOD CONCLUSIONS

We co-located field method testing sites so that results from different methods could be compared. We concluded the following about the various field investigation methods tested:

- Surface Geophysical surveys are a reliable means to characterize large areas of the subsurface, but should be locally calibrated and verified by exploratory boreholes. The geophysical testing results can provide valuable information for siting exploratory boreholes.
- The exploratory borehole lithology observations were the most detailed and accurate characterizations of subsurface materials, and were used as the baseline data to compare other results. The sonic drilling method provided high quality, continuous soil core samples necessary for characterizing the high degree of vertical heterogeneity of the lithology.
- FDEM interpreted results closely matched the exploratory borehole observed net sand thickness in the top 30 feet of the subsurface. FDEM results provide a reliable, relative characterization of the bulk grain size in the top 30 feet of unsaturated alluvial deposits.
- TDEM results did not have the resolution necessary to provide useful subsurface characterization for this project. The method could be adapted to have multiple passes with different equipment configurations (coil sizes). With proper configuration, TDEM may provide a reliable, relative characterization of the grain size and layering in unsaturated alluvial deposits.
- MASW interpreted results closely matched the borehole aquifer cementation and bedrock depth observations. MASW results provided a relative measure of cementation and bedrock depth in unsaturated alluvial deposits.
- ERT results were unreliable and are not recommended for future investigations.
- The interpretations of saturated hydraulic conductivity from the surface and downhole infiltration tests are known to be uncertain. Therefore the interpreted values are meant to give a relative, qualitative permeability estimate and not a precise value suitable for engineering design purposes.
- Surface double-ring infiltrometer testing provided a direct measurement of infiltration rates into surficial soils.
- The methods presented here are not exhaustive, and additional field investigation technologies and methods should continue to be considered.

7.2 LOST CREEK BASIN HYDROGEOLOGIC CHARACTERIZATION CONCLUSIONS

We have made the following conclusions regarding the hydrogeologic characteristics of Lost Creek Basin as they relate to underground water storage planning:



- The alluvial deposits are vertically variable with frequent very thin to thick alternating beds of sand, clay, clayey sand, and sandy clay with some gravelly areas. Intermediate clay beds were observed in all three test holes. These intermediate clay beds have the potential to represent challenges to groundwater recharge. The lateral continuity of these beds and their potential effect on infiltration rates is important and should be better characterized.
- Near surface (shallower than 10 feet below ground surface) fine grained material may inhibit infiltration and should be specifically investigated, even in areas where the deeper deposits are permeable.
- Significant cementation of the alluvial aquifer vadose zone material was observed in two of three test holes.
- The degree of cementation influences permeability and porosity of the alluvial deposits and is a key planning consideration for underground storage at this location.
- Older alluvial deposits generally form topographic highs (terraces), are weakly to well cemented, and display a low to moderate permeability that is less conducive to infiltration. In these older alluvial deposits, the permeability of the subsurface material is not correlated to the gradation and is controlled by the cementation.
- Younger alluvium exists in topographically low spots, was generally loose, poorly cemented, and displayed moderate to high permeability; which is more conducive to infiltration.
- These conclusions are specific to the areas of the Lost Creek Basin that were investigated but are likely applicable to underground storage planning in other alluvial aquifers of the South Platte River and its tributaries.

7.3 RECOMMENDATIONS

Planning of underground storage will benefit from a phased feasibility investigation approach starting with regional or basin-scale investigations followed by localized field investigation. We recommend the following for future feasibility studies of underground storage in the alluvial aquifers of the South Platte River and its tributaries.

7.3.1 Basin Scale or Regional Investigations

- The cementation of the older alluvial deposits is a critical underground storage consideration and is not identified in regional studies, such as the 2011 Lost Creek Basin Aquifer Recharge and Storage Study (CGS, 2011). The reduction in pore space and permeability due to cementation may result in a lower amount of available storage than previously characterized. A desktop study should be conducted to re-evaluate available storage and permeability in the aquifer assuming that the older alluvial deposits are moderately to well cemented.
- Field mapping should be used to refine the available surficial geology and soil maps with a focus on mapping the younger and older alluvial deposits.



- Two-dimensional seismic profiles may help to characterize the distribution of cemented aquifer materials over larger areas.
- Basin/regional scale investigations and underground storage planning may also benefit from three-dimensional geologic modeling and groundwater flow modeling.
- If the conclusions of this study continue to hold true with additional investigation, local field investigations for recharge potential should be focused in the younger Piney Creek and Post Piney Creek alluvial deposits.

7.3.2 Local or Parcel Scale Investigations:

- The potential effect of shallow clay layers in the younger alluvium and the lateral continuity of these clay layers should be further investigated.
- Future investigations should include longer term infiltration testing using test pits or pilot infiltration basins to better characterize realistic infiltration rates. These tests should include observations of the wetting front and changes in groundwater levels throughout the test.
- Future parcel scale investigations should utilize FDEM geophysical surveys to optimize the location of infiltration basins and to target the relatively coarse materials in the top 30 feet of the alluvium.
- Future investigation utilizing the TDEM geophysical method should modify the equipment configuration to investigate the shallower materials in greater resolution.
- Future site investigations should include characterization of alluvium cementation. We recommend the use of seismic geophysical methods (such as MASW) in combination with test holes to characterize the degree of cementation in the alluvium.
- Future local-scale investigations should consider additional methods for investigating the existence and impact of vertical layering and lateral continuity of layers.
- Three dimensional geologic modeling and vadose zone/groundwater modeling should be conducted to evaluate the recharge and associated recovery of recharge water

The results of this study continue to show that there is good potential for underground storage of water in the alluvial aquifers of the South Platte River and its tributaries. Using the recommended investigation methods presented herein will increase the success of underground storage projects.



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| | Bo | orehole 1 | |] | | Bor |
|---------|------|-----------|--------|---|---------|-------|
| Depth F | Rang | ge (feet) | Vs | | Depth F | Range |
| 0.0 | - | 0.3 | 655.4 | 1 | 0.0 | - |
| 0.3 | - | 0.7 | 656.9 | | 0.4 | - |
| 0.7 | - | 1.1 | 650.3 | | 0.9 | - |
| 1.1 | 127 | 1.7 | 644.9 | 1 | 1.5 | - |
| 1.7 | - | 2.4 | 652.0 | | 2.2 | |
| 2.4 | - | 3.3 | 633.4 | 1 | 3.2 | - |
| 3.3 | - | 4.4 | 604.2 | | 4.3 | - |
| 4.4 | - | 5.7 | 561.9 | 1 | 5.8 | - |
| 5.7 | - | 7.5 | 509.0 | | 7.6 | - |
| 7.5 | - | 9.6 | 516.2 | | 9.9 | - |
| 9.6 | - 20 | 12.3 | 672.7 | 1 | 12.8 | - |
| 12.3 | - | 15.7 | 873.6 | 1 | 16.4 | - |
| 15.7 | - | 19.9 | 952.8 | 1 | 20.9 | - |
| 19.9 | - | 25.1 | 848.4 | 1 | 26.5 | - |
| 25.1 | - | 31.7 | 735.4 | | 33.5 | - |
| 31.7 | - | 39.9 | 890.4 | 1 | 42.3 | - |
| 39.9 | - | 50.2 | 1116.1 | 1 | 53.2 | - |
| 50.2 | - | 58.3 | 1071.7 | 1 | 57.8 | - |
| 58.3 | - | 69.9 | 1006.1 | 1 | 69.4 | - |
| 69.9 | - | 81.6 | 947.5 | 1 | 81.0 | - |
| 81.6 | - | 93.2 | 870.1 | 1 | 92.5 | - |
| 93.2 | - | 104.9 | 1227.7 | | 104.1 | - |
| 104.9 | - | 116.5 | 1679.0 |] | 115.7 | - |
| 116.5 | - | 128.2 | 2054.6 |] | 127.2 | - |
| 128.2 | - | 139.8 | 2255.5 |] | 138.8 | - |
| 139.8 | - | 151.5 | 2275.2 |] | 150.3 | - |
| 151.5 | - | 163.1 | 2235.1 | | 161.9 | - |
| 163.1 | | 174.8 | 2014.8 | | 173.5 | - |
| 174.8 | - | 186.4 | 1799.9 | | 185.0 | - |
| 186.4 | - | 198.1 | 1576.9 | | 196.6 | |
| 198.1 | - | 209.8 | 1220.1 | | 208.2 | - |
| 209.8 | 4 | 221.4 | 1080.8 | | 219.7 | - |
| 221.4 | - | 276.8 | 2952.2 | | 2.5 | |

| Borehole 4 | 1 | |
|--------------|--------|----|
| Range (feet) | Vs | De |
| - 0.4 | 513.8 | 0 |
| - 0.9 | 569.3 | |
| - 1.5 | 639.7 | 0 |
| - 2.2 | 598.4 | 1 |
| - 3.2 | 668.7 | 2 |
| - 4.3 | 592.7 | 2 |
| - 5.8 | 511.9 | 3 |
| - 7.6 | 450.9 | 5 |
| - 9.9 | 524.3 | 6 |
| - 12.8 | 699.0 | 8 |
| - 16.4 | 656.0 | 1 |
| - 20.9 | 671.7 | 1 |
| - 26.5 | 863.4 | 1 |
| - 33.5 | 1003.2 | 2 |
| - 42.3 | 834.2 | 2 |
| - 53.2 | 879.3 | 3 |
| - 57.8 | 1125.5 | 5 |
| - 69.4 | 905.3 | 6 |
| - 81.0 | 757.7 | 6 |
| - 92.5 | 1066.4 | 7 |
| - 104.1 | 1450.0 | 8 |
| - 115.7 | 1706.9 | 9 |
| - 127.2 | 1835.9 | 10 |
| - 138.8 | 1833.8 | 11 |
| - 150.3 | 1760.1 | 12 |
| - 161.9 | 1665.2 | 12 |
| - 173.5 | 1554.4 | 13 |
| - 185.0 | 1468.1 | 14 |
| - 196.6 | 1421.6 | 15 |
| - 208.2 | 1407.1 | 16 |
| - 219.7 | 1418.7 | |
| - 274.7 | 2327.3 |] |

DATE: 6/20/2017

Provided by Olson Engineering

Figure 14 Multi-channel Analysis of Surface Waves (MASW) Results

1221 Auraria Parkway | Denver, CO 80204 303-455-9589 | 800-453-9589 | Fax: 303-455-0115 www.LREwater.com

Appendix A – Lost Creek Basin Geology Map

MISCELLANEOUS INVESTIGATIONS SERIES

intrusive bodies in the amphibolite. In the Montezuma region near the west edge of

he quadrangle this unit appears to represent a pre-metamorphic intrusive

complex. May contain some layers of calc-silicate gneiss near contact with

containing beds of quartzite-pebble conglomerate and beds of biotite sillimanite

schist. Occurs only in Coal Creek syncline in the northeastern part of Jefferson

unlayered amphibolite, but contains a minor amount of layered hornblende gneiss,

felsic gneiss, and calc-silicate gneiss. Probably represents basaltic or andesitic lava

that contains numerous layers of calc-silicate gneiss and calc-silicate quartzite,

some layers of felsic gneiss, muscovite quartzite, and, in the eastern part of the

ront Range, a few layers of marble. Probably derived from interbedded sediment

MPHIBOLITE AND CALC-SILICATE GNEISS (PRECAMBRIAN X)-Amphibolite

UARTZITE (PRECAMBRIAN X)—Gray, coarse-grained micaceous quartzite

HORNBLENDE GNEISS AND AMPHIBOLITE (PRECAMBRIAN X)-Predominantly

ows, plutons, and minor interbedded tuff and sediment

mphibolite and calc-silicate gneiss (Xhs)

Reinecker Ridge Volcanic Member and lower conglomerate and arkose DAWSON ARKOSE (EOCENE, PALEOCENE, AND UPPER CRETACEOUS) Siliceous conglomerate and conglomeratic sandstone facies-Yellowish-gray to rayish-brown arkosic conglomerate and coarse-grained to conglomeratic arkosic andstone. Matrix is commonly hard, siliceous, and may be opalized or tuffaceous. Locally contains tuffite lenses which are very hard and break with conchoidal racture across sand grains. Thickness is as much as 30 m Arkosic sandstone and claystone facies-Yellowish-gray to light-gray, commonly ironstained, coarse- to fine-grained, locally conglomeratic, massive to crossbedded

of the 600,000 year old Pearlette type

and in the upper 120 m contains hard, grav to vellowish-grav tuffite beds that brea with conchoidal fracture. A weathered crystal tuff occurs in the lower part of th facies in the central part of the map area southeast of Denver. Biotite from this tu has been dated at 56.5±1.9 m.y. (R. F. Marvin, written commun., 1975); this, and a fission-track age determination of 54.8±5.6 m.y. on zircon by C. W. Naese (unpub. data, 1977) from what is believed to be the same tuff, indicate that the tu is of late Paleocene or early Eocene age. Thickness more than 305 m in surface exposures; estimated to be about 915 m in subsurface (Kittleman, 1956) but this igure includes about 210 m of carbonaceous shale and coal facies that is include in unit TKdI on this map. In cross-section Td includes Tdb Fine-grained sandstone, carbonaceous shale and lignite facies-Light-gray, hard, limy, fine-grained, concretionary sandstone and soft, rusty, fine-grained sandsto beds containing limy "cannonball" concretions with brown rinds interbedded

gray to dark-brown and black carbonaceous shale and clay and numerous lignit eds. Small spherical ironstone concretions with coal nucleii are present locally, and large, irregularly-shaped ironstone crusts and lenses are common in soft rust sandstone layers. A hard gray tuffite bed several cm to 30 cm thick is present at o near the top of the facies throughout much of the southeast part of the map area. A onspicuous clinker horizon occurs near the top of the facies in the southeast pa he map. The clinker bed is overlain by the tuffite in at least one locality. The top of the facies is also commonly marked by a residue of quartz and chert pebbles and occasionally by a red, yellow, and white variegated clay paleosol zone. The clay one commonly contains abundant quartz and some feldspar grains and locally

FELSIC GNEISS (PRECAMBRIAN X)-Biotite-quartz-plagioclase-microcline gneiss Lyons Sandstone-Light-gray to grayish-orange, crossbedded, fine- to containing various proportions of plagioclase and microcline; contains layers and medium-grained quartzose sandstone and some conglomerate, siltstone, and lenses of amphibolite and, less commonly, layers of biotite-quartz-plagioclase mudstone. Thickness 46-76 m PIPF FOUNTAIN FORMATION (PERMIAN AND PENNSYLVANIAN)-Grayish-red, schist. Grain-size ranges from fine to coarse. In areas west and south of Mt. Evans, reddish-brown, moderate-red, and gray coarse-grained, arkosic sandstone with he amphibolite occurs as small irregular bodies that appear intrusive into the felsic gneiss. May represent flows, pyroclastics, and hypabysal intrusive rock. Interfingers lenses of siltstone and fine-grained sandstone, locally conglomeratic and with and grades into biotite gneiss (Xb) west of Golden AYERED GNEISS (PRECAMBRIAN X)—Biotite-quartz-plagioclase schist and gneis crossbedded. Unit forms prominent hogbacks along east edge of Front Range and underlies Manitou Park (T. 10-12 S., R. 69 W.). Thickness 240-900 m; about 1,200 m in subsurface near south edge of quadrangle MAROON FORMATION (LOWER PERMIAN AND UPPER AND MIDDLE felsic gneiss, amphibolite, calc-silicate gneiss, and impure quartzite. Probably represents interlayered and intergrading tuff, tuffaceous sediment and sediment BIOTITE GNEISS (PRECAMBRIAN X)—Biotite-quartz-plagioclase schist and gneiss, PENNSYLVANIAN)-Grayish-red to moderate-red and gray arkose, siltstone, commonly contains abundant sillimanite and less abundant muscovite. Some and mudstone in western South Park. Bleached to green and gray colors in the avers of cordierite-biotite gneiss and of garnet-biotite gneiss. Locally a few layers of arryall Creek (T. 8 S., R. 77 W.) area by contact metamorphism. About 1,700 m omblende gneiss and calc-silicate gneiss. Lenses, pods, and thin layers of Pm MINTURN FORMATION (MIDDLE PENNSYLVANIAN)-Grayish-red, moderpegmatite abundant. In some regions layers and lenses of granodiorite and qui nonzonite are also abundant and rock grades to migmatite. Probably derived from ate-red, reddish-brown, and gray arkosic siltstone and mudstone containing beds shale, siltstone, and sandstone of gray limestone. At least 1,100 m thick CORDIERITE-BEARING AND GARNET-BEARING SILLIMANITE-BIOTITE aporite-bearing facies of Minturn Formation-Gray shale, limestone, siltstone, GNEISS (PRECAMBRIAN X)-Dark-gray, fine- to medium-grained gneiss. sandstone, white silty limestone, anhydrite, and gypsum in southwest corner of Mapped in north-central and east-central Front Range. Gradational contacts with PPmb MAROON FORMATION, MINTURN FORMATION AND BELDEN (MIDDLE AND niotite gneiss. Probably derived from iron and magnesium-rich graywacke and LOWER PENNSYLVANIAN) FORMATION—Shown only in cross section; may p€u UNDIFFERENTIATED PRECAMBRIAN ROCKS-Shown only in cross sections contain some older Paleozoic rocks at west part of B-B'.

DAKOTA SANDSTONE, MORRISON FORMATION, AND ENTRADA SAND-

NIOBRARA FORMATION, CARLILE SHALE, GREENHORN LIMESTONE

SANDSTONE (PERMIAN)-Shown along east side of the Front Range

MORRISON AND RALSTON CREEK FORMATIONS (UPPER JURASSIC), LYKINS

central margin of mapped area

FORMATION—Shown only in cross section

varicolored chalcedony. Thickness 18-46 m

STONE—Unit includes about 45 m of Chinle Formation of Late Triassic age south

of Boreas Pass (T. 7 S., R. 77 W.) (Poole and Stewart, 1964). Shown near west

GRANEROS SHALE, DAKOTA SANDSTONE, AND MORRISON

FORMATION (LOWER TRIASSIC AND UPPER PERMIAN), AND LYONS

Ralston Creek Formation-Varicolored claystone interbedded with some lime-

stone, siltstone, and sandstone. Contains thin beds and disseminated nodules of

Lykins Formation—Gravish-red shaly mudstone with thin beds of light-gray, very

fine-grained sandstone and some "crinkled" limestone in lower part. Thickness

d, Dawson Formation (arkosic sandstone facies) Pre-Laramie rocks

> Fence diagram showing relations of Upper Cretaceous and Tertiary rock units above the base of the Laramie Formation in the Denver Basin Dashing shows natural profile on top of thick Quaternary deposit in SE corner of diagram. (M-Q and Q-S)]

T.12 S.

S" RESTON, VA MAY 0 9 199 BRA

M(200) no.1163 sheet c,2

For sale by Branch of Distribution, U.S. Geological Survey, Box 25286, Federal Center, Denver, CO 80225

Appendix B – Test Hole Summary Logs

O-1 Total depth: 80 UTM Easting (ft): UTM Northing (ft): Elevation (ft):

Date drilled: 4/19/18 Permit number: N/A Project number: 1412AUW02 Field Geologist: Joel Barber

Leonard Rice ENGINEERS, INC.

1221 Auronia Parkway | Denver, CO 80204 303-455-9589 | 800-453-9589 | Fax: 303-455-0115 www.LREwater.com

| Depth (ft) | | Lithology Log | | | | Resis | tivit | ity MASW Vs (ft/s) | | | | | | | Permeability | | | | | | |
|------------|-------|--|-----|---|------------------|-------|-------|-----------------------|-----|--|-----|---|------|------|--------------|--|---|-------|----|---|----|
| 0 | | CL: Clay | | | | | | | | | | , | | | | | | (100) | | | |
| - | | SP: Poorly Graded Sand. Tan. Dry. Weak cohesion in sample. | | | | | | | | | | | | | | | | | | | |
| - | | | | | | | | | | | | | | | | | | | | | |
| | | | | | | | | | | | | | | | | | | | | | |
| - | | | _ | | | | | | | | | | | | | | | | | | |
| - | | SP-SC: Poorly Graded Sand w/ Clay. Brown mottled with white. Dry. Calcite precipitate. | - | | | | | | | | | | | | | | | | | | |
| -10 | •••• | SP: Poorly Graded Sand. Tan. Dry. Weak cohesion in sample. | - | | | | | | | | | | | | | | | | | | |
| | | SC: Clayey Sand. Brown mottled w/ white. Cemented. Difficult drilling. | | | | | | | | | | | | | | | | | | | |
| | | | | | | | | | | | | | | | | | | | | | |
| - | | | | | | | | | | | | | | | | | | | | | |
| - | | SW: Well Graded Sand. Tan. Dry. Weakly cemented. | | | | | | \mathbf{N} | | | | | | | | | | | | | |
| - | | | _ | | | | | | | | | | | | | | | | | | |
| 20 | | SC: Clayey Sand. Brown mottled w/ white. Cemented. Difficult drilling. | | | | | | | | | | | | | | | | | | | |
| -20 - | | SW: Well Graded Sand. Tan. Dry. Weakly cemented. | | | | | | | | | | | | | | | | | | | |
| - | | | | | | | | | | | | | | | | | | | | | |
| - | | CL: Clay. Light brown. Dry. Calcite. Very stiff, apparent strong cementation. | | | | | | | | | | | | | | | | | | | |
| - | | | 0.1 | | 1 | | 10 | 1 | 00 | | | | | | 0.1 | | 1 | | 10 |) | 1(|
| | | CL: Sandy Clay. Brown. Dry. Very weak cementation. | | | | | | | | | | | | | | | | | | | |
| | | | | | | | | / | | | | | | | | | | | | | |
| -30 - | | SW: Well Graded Sand. Tan. Dry. Weakly to moderately cemented. | | | | | | | 400 | | 800 | | 1200 | 160' | 0 | | | | | | |
| - | | | | | | | / | | | | | | | | | | | | | | |
| - | | | | | | / | | | | | | | | | | | | | | | |
| _ | | | - | | | / | | | | | | | | | | | | | | | |
| | | SP: Poorly Graded Sand with Clay Beds. Tan. Dry. weakly cemented. | | | | / | | | | | | | | | | | | | | | |
| - | | CL: Sandy Clay. Brown. Dry. | | | $\left \right $ | | | | | | | | | | | | | | | | |
| -40 - | | SW-SC: Well Graded Sand w/ Clay. Brown. Moist. Moderate dry strength from cohesion or cementation. | | / | | | | | | | | | | | | | | | | | |
| - | | | | | | | | | | | | | | | | | | | | | |
| - | | | | / | | | | | | | | | | | | | | | | | |
| | | CL: Sandy Clay. Brown. Dry. Very high dry strength. Sonic rig cored through sample. | 0.1 | | 1 | | 10 | 1 | 00 | | | | | | 0.1 | | 1 | | 1(|) | 1(|
| - | | SP: Poorly Graded Sand. Tan. Dry. Very weak cohesion/cementation. | | | | | | | | | | | | | | | | | | | |
| - | | | | | | | | | | | | | | | | | | | | | |
| -50 — | | SP: Calcite rich Poorly graded Sand. White. Dry. Calcite "powder" | | | | | | | | | | | | | | | | | | | |
| - | | SP: Poorly Graded Sand. Tan. Dry. Very weak cohesion/cementation. | | | | | | | | | | | | | | | | | | | |
| | ••••• | SW: Gravelly Well Graded Sand. Tan. Dry. No apparent cementation. | | | | | | | | | | | | | | | | | | | |
| | | | | | | | | | 400 | | 800 | | 1200 | 160 | 10 | | | | | | |
| - | | CL: Sandy Clay. Brown. Dry. Very stiff/well cemented. Sample can't be broken by hand. | | | | | | | | | | | | | | | | | | | |
| - | | | | | | | | | | | | | | | | | | | | | |
| -60 - | | SP: Poorly Graded Sand with Clay Laminations. Horizontal sructure - Laminated to thinly bedded. Weakly cemented. | 1 | | | | | | | | | | | | | | | | | | |

0-4

Total depth: 60 UTM Easting (ft): UTM Northing (ft): Elevation (ft): Date drilled: 4/19/18 Permit number: N/A Project number: 1412AUW02 Field Geologist: Joel Barber

ENGINEERS, INC. 1221 Aurania Parkway | Denver, CO 80204 303-455-9589 | 800-453-9589 | Fax: 303-455-0115

303-455-9589 | 800-453-9589 | Fax: 303-455-0115 www.LREwater.com

| Depth (ft) | Lithology Log | | TDEM Resistivity (Ohm-m) | | | | | | MASW Vs (ft/s) | | | | | Permeability | | | | | |
|------------|--|-----|-----------------------------|---|---|-----------|---|-----|-------------------|----------|-----|------|-----|--------------|------|--|-----|---|---|
| 0 | SC: Clayey Sand. Brown. Moist. Organics. | | | | | <u></u> , | | | 5 | | | , | | | | | / | | |
| - | SP-SC: Poorly Graded Sand s/ Some Clay. Light Brown. Dry. | | | | | | | | | <u> </u> | | | | | | | | | |
| - | | | | | | | | | | | | | | | | | | | |
| - | | | | | | | | | | | | | | | | | | | |
| - | SC: Clayey Poorly Graded Sand. Light Brown. Dry. Weakly cemented, low dry strength. | | | | | | | | | | | | | | | | | | |
| -10 | | | | | | | | | | | | | | | | | | | |
| | | | | | | | | | | | | | | | | | | | |
| - | | | | | | | | | | | | | | | | | | | |
| - | CL: Sandy Clay. Light brown. Dry. Strong dry strength/cementation. "Cored" by sonic rig. | _/ | | | | | | | | | | | | | | | | | |
| - | | | | | | | | | | ļ | | | | | | | | | |
| - | | | | | | | | | | | | | | | | | | | |
| -20 — | CL: Sandy Clay. Light brown w/ white staining. Dry. Strong dry strength/cementation. "Cored" by sonic rig. | | | | | | | | | | | | | | | | | | |
| _ | SP-SC: Poorly Graded Sand w/ Clay. Weak cementation/dry strength. | | | | | | | | | | | | | | | | | | |
| _ | SW: Gravelly Well Graded Sand. Tan. Dry. No apparent cementation. | | | | | | | | | | | | | | | | | | |
| | SW-SC: Well Graded Sand w/ Clay. Brown. Dry. Moderately Cemented. | 0.1 | | 1 | | 1(|) | 100 |) | | | | | 0.001 | 0.01 | | 0.1 | 1 | 1 |
| - | SW: Gravelly Well Graded Sand. Tan. Dry. Well cemented. | | | | | | | | | | | | | | | | | | |
| - | SW-SC: Well Graded Sand w/ Clay. Brown. Dry. Moderately Cemented. | | | | | | | / | | | | | | | | | | | |
| -30 — | SW: Gravelly Well Graded Sand. Tan. Dry. No apparent cementation. | | | | | | | / | 400 | ç | 100 | 1200 | 160 | 10 | , | | | | |
| _ | CL-SC-SP: Interbeded CL, SC, SP-SC, and SP. Brown. Dry. Well cemented with moderate to high dry strength. Predominately sand with thin clay beds. Horizontal bedding structure. | | | | | | | | 100 | | | | | | | | | | |
| | | | | | | | / | | | | | | | | | | | | |
| - | | | | | | / | | | | | | | | | | | | | |
| - | | | | | | / | | | | | | | | | | | | | |
| - | | | | | / | | | | | | | | | | | | | | |
| -40 | | | | | / | | | | | | | | | | | | | | |
| - | SW: Well Graded Sand. Tan. Dry. Well cemented with a weak horizontal structure. Contains infreqent very thin to | | | / | | | | | | | | | | | | | | | |
| _ | thin clay beds. | | / | | | | | | | | | | | | | | | | |
| | | 0.1 | / | 1 | | 1(|) | 100 |) | | | | | 0.001 | 0.01 | | 0.1 | 1 | 1 |
| | | | | | | | | | | | | | | | | | | | |
| - | | | | | | | | | | | | | | | | | | | |
| -50 — | | | | | | | | | | | | | | | | | | | |
| - | | | | | | | | | | | | | | | | | | | |
| - | | | | | | | | | | | | | | | | | | | |
| - | | | | | | | | | 400 | 8 | 00 | 1200 | 160 | 0 | | | | | |
| | SW: Well Graded Sand. White. Dry. Calcite cement broken up into powder by sonic rig. | | | | | | | | | | | | | | | | | | |
| | SW: Well Graded Sand. Tan. Dry. Very well cemented, appears as sandstone with calcite cement. | | | | | | | | | | | | | | | | | | |
| -60 | •• | _ | | | | | | | | | | | | | | | | | |

O-5

Total depth: 85 UTM Easting (ft): UTM Northing (ft): Elevation (ft): Date drilled: 4/23 to 4/24/18 Permit number: N/A Project number: 1412AUW02 Field Geologist: Joel Barber

1221 Auraria Parkway | Denver, CO 80204 303-455-9589 | 800-453-9589 | Fax: 303-455-0115 www.LREwater.com

| Depth (ft) | Lithology Log | TD | EM Resistivity (Ohm-m) | MASW Vs (ft/s) | Permeability (ft/d) |
|------------|--|----------------------|---------------------------|-------------------|------------------------|
| 0 | SC: Clayey Poorly Graded Sand with Organics. Dark Brown. Moist. Contains frequent roots. | • | | | |
| - | SC: Clayey Poorly Graded Sand. Brown. Moist. Some white (calcite) staining. | | | | |
| - | | | | | |
| | | | | | |
| - | | | | | |
| - | CL: Sandy Clay. Dark brown. Moist. | | | | |
| -10 | SC: Clayey Poorly Graded Sand. Brown. Moist. Some white (calcite) staining. | | | | |
| | | | | | |
| - | SP-SC: Poorly Graded Sand w/ Clay. | | | | |
| - | | | | | |
| - | | | | | |
| | CL: Sandy Clay. Brown. Moist. | | | | |
| - | SW: Well Graded Sand w/ Some Gravel. Tan. Moist. | | | | |
| -20 - | SC: Clayey Sand. Brown to tan. Moist. | | | | |
| - | | | | | |
| | SP: Poorly Graded Sand. Tan. Moist. | | | | |
| - | | | | | |
| - | | 0.1 | 1 10 10 | | 0.1 |
| - | SC: Clavey Boorly Graded Sand Tap. Moint to wat | | | | |
| | SP: Poorly Graded Sand. Tap. Moist | | | | |
| -30 - | Sr. Fooling Graded Salid. Tail. MOISL. | | | 400 800 1200 | 1600 |
| - | SP: Poorly Graded Sand. Tan. Moist | | | | |
| | | | | | |
| | SP-SC: Poorly Graded Sand w/ Clay. Tan. Moist. Weakly cemented. Horizontal thin bedding structure. | | | | |
| - | SW: Gravelly Well Graded Sand. Light brown and mottled. Moist. Very weakly cemented. | | | | |
| - | | | | | |
| 40 | SP-SC: Poorly Graded Sand w/ Clay. Tan. Moist. Weakly cemented. Horizontal thin bedding structure. | | | | |
| -40 | | | | | |
| - | | | | | |
| - | | / | | | |
| | SP: Poorly Graded Sand. White. Dry. Broken up calcite powder in sample. | 0.1 | 1 10 10 | | 0.1 1 10 100 1000 |
| | SP-SC: Poorly Graded Sand w/ Clay. Tan. Moist. Weakly cemented. Horizontal thin bedding structure. | | | | |
| - | SP: Poorly Graded Sand. Light brown. Moist. | | | | |
| -50 - | SW: Well Graded Sand. Light brown. Moist. | | | | |
| | SP-SC: Poorly Graded Sand w/ Clay. Tan. Moist. Weakly cemented. Horizontal thin bedding structure. | | | | |
| | SW: Well Graded Sand. Light brown. Moist. | | | | |
| - | CL: Clay w/ Sand. Brown. Moist. Sonic barel "cored" throug clay. | | | | |
| - | SC: Well Graded sand w/ Clay. I an mottled with dark brown. Moist. | | | 400 800 1200 | 1600 |
| | SP-SC: Poorly Graded Sand W/ Clay. Tan. Moist. Weakly cemented. Horizontal thin bedding structure. | | | | |
| | SP: Poorly Graded Sand. Tan. Dry. Weakly cemented. | | | | |
| -60 | | | | | |
| - | SW: Interbedded Well Graded Sand with varying colors from black to tan to brown. Moist. | | | | |
| _ | | | | | |
| | SC: Clayey poorly graded sand. Light brown. Moist. | 0.1 | 1 10 10 | | 0.1 1 10 100 |
| - | Calcite staining. | | | | |
| - | SW: Well Graded Sand. Brown. Moist to wet. | | | | |
| -70 | | | | | |
| | | | | | |
| - | | | | | |
| - | | | | | |
| | | | | | |
| | | | | | |
| - | CL: Clay. Grey. Wet. | | | | |
| -80 - | SC: Clayey Sand. Tan. Wet. | | | 400 800 1200 | 1600 |
| | SW: Well Graded Sand. Brown. Moist to wet. | | | | |
| - | Denver: Denver Formation. Blue shale and very fine grained sandstone. | | | | |
| - | | | | | |
| I | 1 of 2 | J | | | |

Leonard Rice **O-5** Total depth: 85 Date drilled: 4/23 to 4/24/18 UTM Easting (ft): -104.517251 Permit number: N/A Project number: 1412AUW02 1221 Auraria Parkway | Denver, CO 80204 303-455-9589 | 800-453-9589 | Fax: 303-455-0115 UTM Northing (ft): 39.880174 Field Geologist: Joel Barber Elevation (ft): 5246 www.LREwater.com **TDEM Resistivity** MASW Vs (ft/s) Permeability (ft/d) Depth (ft) Lithology Log (Ohm-m) 00 00 -90 -100 -800 1200 16000.1 100 400 00 -110 -120 --130 800 1200 1600 400 -140 -

| - | 0.1 | 10 100 1000 |
|------|-----|-------------|
| -150 | | |